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Title: The 4.2 ka event: A review of palaeoclimate literature and directions for future research

Year: 2024

Version: Final draft

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Please cite the original version:

Helama, S. (2024). The 4.2 ka event: A review of palaeoclimate literature and directions for future research. *The Holocene*, 34(9), 1408-1415. <https://doi.org/10.1177/09596836241254486>

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The 4.2 ka event: a review of palaeoclimate literature and directions for future research

Journal:	<i>The Holocene</i>
Manuscript ID	HOL-23-0178.R1
Manuscript Type:	Forum Article
Date Submitted by the Author:	n/a
Complete List of Authors:	Helama, Samuli; Natural Resources Institute Finland,
Keywords:	4.2 ka event, Holocene, palaeoclimatology, chronostratigraphy, ice-rafted debris events, monsoon
Abstract:	<p>In recent years, much evidence has been presented on the 4.2 ka event. A review of 317 palaeoclimate papers shows that dry conditions were common during the event, especially from Eastern Mediterranean to India. The 4.2 ka event was not, however, a global drought event. Wet conditions were reported especially for central/northern Europe and sub-Saharan Africa. The 4.2 ka event is typically characterized either as short (4.2-4.0 ka) or long (4.4-3.8 ka) episode, possibly developing over an extended interval of time, in keeping with the North Atlantic forcing and correlating with the Bond 3 event of ice-rafted debris. This forcing is understood to drive a southward migration of the Inter-Tropical Convergence Zone (ITCZ), resulting in decreased rainfall over most of the Asian monsoon region, with possibility that an interplay of El Niño–Southern Oscillation (ENSO) and North Atlantic Oscillation (NAO) has modulated the global circulation. Cold conditions were also reported but less frequently, in comparison to the 8.2 ka event, Dark Ages Cold Period and Little Ice Age. Some high-resolution records show a double peak structure of which two anomalies are tree-ring dated to 4.14-4.05 ka and 3.97 ka. Accurately and precisely dated high-resolution records indicative of various climatic variables, especially outside of the traditional study region (Mediterranean–Middle East–India–China), including reconstructions of the ENSO and NAO histories and ITCZ migrations, are crucially needed for rigorous examination of the global scale characteristics of the 4.2 ka event and its forcings. Such research seems to be just beginning.</p>

Introduction

Holocene climate history is characterized by distinct climate variability at the millennial and centennial timescales (Mayewski et al., 2004; Wanner et al., 2008, 2011; Solomina et al., 2015). One of the frequently cited climatic episodes dates back around four millennia and is commonly termed the 4.2 ka event or 4.2 ka BP event (Rousseau et al., 2019). Research of this event has its own history that goes back decades. Early studies recognised reductions in monsoon precipitation and coinciding weakening of the Indus Valley civilisation (Singh, 1971). Synchronous increases in aridity (Figure 1a) around same time in the Mesopotamia, Aegean, Egypt, Levant and the Indian subcontinent are generally understood to have driven the documented collapses of the ancient rain-fed agriculture civilisations (Weiss et al., 1993; Cullen et al., 2000; deMenocal, 2001). The same region extended to China, climate-culture nexus and drastic changes in monsoon circulation have all remained in the core of the multidisciplinary research focussing the 4.2 ka event since those days (Ran and Chen, 2019). In this context, a wider spread of the coinciding droughts and apparent global linkages have been illustrated also for North American sites (Booth et al., 2005).

More recently, an isotopic ($\delta^{18}\text{O}$) excursion recorded in Mawmluh Cave speleothem in northeast India (Figure 1), reflecting the reduction in monsoon rainfall during the 4.2 ka event (Berkelhammer et al., 2012), was formally ratified as Global Boundary Stratotype Section and Point (GSSP) for the lower boundary for the Meghalayan Stage/Age, and accordingly the boundary between the Middle and Late Holocene, by the Executive Committee of the International Union of Geological Sciences (Walker et al., 2018, 2019a, 2019b). The ratification was supported, among other evidence, by a dataset that had been collated by Mayewski et al. (2004), who had described synchronous rapid climatic disruptions with wide spatial distribution between 4.2 and 3.8 ka, in connection with major atmospheric circulation changes typified as a “cool poles, dry tropics” pattern. Apart from droughts, however, wet conditions seem to characterize the 4.2 ka event at high latitudes, including the

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3 Mount Logan ice-core succession that constitutes an auxiliary stratotype for the Mawmluh Cave
4 (Figure 1b) (Fisher et al., 2008; Walker et al., 2019b), as well as in parts of sub-Saharan Africa (Booth
5 et al., 2005; Walker et al., 2012; Railsback et al., 2018, 2022). Moreover, advances of Northern
6
7 Hemisphere glaciers over the same period, in keeping with cooling in the North Atlantic sector, have
8
9 been identified (Denton and Karlén, 1973; Solomina et al., 2015). With these regards, the foregoing
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11 papers (deMenocal, 2001; Mayewski et al., 2004; Booth et al., 2005; Solomina et al., 2015), like
12
13 several other authors, have discussed the 4.2 ka event in the context of a North Atlantic events of
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15 ice-rafted debris as recorded by Bond et al. (1997, 2001). The North Atlantic data show a cold-water
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17 pulse commonly referred to as Bond 3 event at around 4.2 ka and suggesting reduced solar radiation
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19 as driver of the Bond events over Holocene time scales (Figure 1c). Putatively, such forcings in
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21 addition to ENSO-driven (El Niño–Southern Oscillation) sea surface temperature patterns (Booth et
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23 al., 2005) could explain the global extent of the event, although the teleconnections involved and
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25 actual physical mechanisms underlying the possible linkages (Gupta et al., 2021; Banerji and
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27 Padmalal, 2022; Renssen, 2022) remain to be further explored.

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36 More critical views of the event have also been expressed. No convincing support for the global
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38 spread of the event was found in a multi-proxy study (Wanner et al., 2008). A global analysis of
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40 speleothem records did not detect notable isotopic excursions around 4.2 ka (Parker and Harrison,
41
42 2022). These authors concluded that the lack of consensus about the event could stem from its
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44 spatial complexity, poor age constraints and/or its weak climatic signals. Similar to criticism on
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46 palaeoclimate indications of the 4.2 ka event, worries about the Middle–Late Holocene boundary,
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48 concerning its dating accuracy together with its' putative global spread, have been created (Voosen,
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50 2018).

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57 Clearly, there seems to be a need for studies that aim to delineate the evidence presented for the
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59 4.2 ka event. This paper constitutes a review of evidence on climate variability during the 4.2 ka
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3 event (Table S1). A literature search for “4.2 ka event” and for its various synonyms (Table S2) was
4
5 performed (Google Scholar (<http://scholar.google.com/>), accessed 26 June 2023) and all papers
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7 citing the term with new proxy indications of climatic or environmental evidence on the event,
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9 published in international English-language peer-reviewed journals, were included in this review.
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11 The number of explored papers totalled 317 and exceeds by ~tenfold the number of papers included
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13 in previous literature collections (Booth et al., 2005; Railsback et al., 2018). Pure literature syntheses
14
15 were not included. Archaeological papers without natural proxy data were excluded to avoid
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17 mismatches arising from differing scales of human responses to any potential climatic driving
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19 mechanisms (Jaffe et al., 2021).
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25 Characterizing the 4.2 ka event literature

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30 It seems the actual phrase “4.2 ka event” was first used by Staubwasser et al. (2003) who discussed
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32 the Holocene drought/monsoon cycles and the demise of the Indus Valley civilization around that
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34 time. The literature that follows constitutes the foundation and rationale for examining the
35
36 characteristics of this event and how it has been detailed using a large set of palaeoclimatic records.
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38 With these regards, there is a notable increase of literature citing the 4.2 ka event since 2018 (Figure
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40 2a). Interestingly, this is the year of the formal ratification of the Meghalayan GSSP (Walker et al.,
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42 2018). It appears plausible that the ratification increased the interest on, or at least awareness of,
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44 climatic/environmental variations pertaining to the 4.2 ka event.
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50 The sites encompassing China, India and the Mediterranean/Levant account for approx. two thirds
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52 of the studies (Figure 2b). Proxy indications were adopted as reported in the reviewed literature. Dry
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54 conditions, recorded as droughts or increases in aridity, were the most common (47% of the studies)
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56 proxy indication (Figure 2c), such conditions dominating (>50% of the indications) especially in India
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58 and Middle East/Eastern Mediterranean (Figure 3). The 4.2 ka event was not, however, a global dry
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3 event but wet conditions (16%) were also reported, especially in central/northern Europe, sub-
4 Saharan Africa, and in parts of China. This means that moist conditions have been reported also from
5 the Asian monsoon region. This would agree with the previously defined tripolar precipitation
6 pattern whereby rainfall decreased over India and southern and northern China and increased over
7 central China (Tan et al., 2018; Zhang et al., 2018; Lin et al., 2022). Cold (19%) conditions were also
8 reported but less frequently. This differentiates the 4.2 ka event from other major Holocene climate
9 episodes, the 8.2 ka event (Alley and Ágústadóttir, 2005; Rohling and Pälike, 2005), Dark Ages Cold
10 Period (Helama et al., 2017, 2021) and Little Ice Age (Lamb 1995; Grove 2004) which were
11 dominated by cold conditions. Clearly, the 4.2 ka event was prominently a hydroclimatic episode and
12 certainly not a warm period almost anywhere (2% of the studies).
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28 Many of the studies were based on multiple proxies (40%) (Figure 2d). Materials were mostly dated
29 either using radiocarbon (^{14}C) determinations (77%), followed by uranium–thorium (14%) and
30 optically stimulated luminescence methods (9%). The starting and ending years and peak/mid-point
31 of the event were adopted as reported in the original studies. The mid-point was on average dated
32 to 4.16 ka. A large number (46) of papers reported the date of the event simply as “4.2 ka”. More
33 detailed chronological examination showed an average starting and ending dates of 4.37 ka and 3.81
34 ka for the event, respectively. Rounded to 100 years, the most common starting years were 4.2 ka
35 (44 studies), 4.3 ka (32) and 4.4 ka (30), the most frequent ending years being 3.8 ka (35), 4.0 ka (31)
36 and 3.9 ka (24) (see Figure 4a).
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50 In the literature, the 4.2 ka event appears either short, with a duration of 200 years, or long, in
51 which case it could span over half-a-millennia. That is, the most usual chronological positions of the
52 event were 4.2-4.0 ka (11 studies), 4.3-3.8 ka (10) and 4.4-3.8 ka (9). The most common durations of
53 the event were 200 (29), 400 (21), and 500 (20) years.
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3 Hydroclimatic events and those recorded in the Asian monsoon region seemed to have both started
4 and ended slightly later than other events (Figure 4b). In the monsoon region, the events had ended
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6 ~110 years later than elsewhere, which could possibly reflect findings that the drought of the Indian
7
8 summer monsoon may have markedly extended over the early Late Holocene (Gupta et al., 2021;
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10 Scroton et al., 2023). Even so, this difference was not found to be statistically significant (t test, $p >$
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12 0.05).
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19 Not all published data appears to concur with the evidence of the 4.2 ka event. These studies do not
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21 seem to be clustered at any one type of proxy or region (e.g. Swindles et al. 2013; Mishra et al.,
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23 2015; Jaouadi et al., 2016; Fritz et al. 2018; Li et al., 2018a; Ön et al., 2018; Hang et a., 2020; Li et al.,
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25 2021; Yang et al., 2021).
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30 Multiple phases within the event

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34 Several papers have illustrated and discussed climate anomalies within the nucleus (4.2-4.0 ka) of
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36 the 4.2 ka event. A double-peak structure has been described at least in the following studies (see
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38 right-hand plots in Figure 1): dry pulses at 4.03 and 3.96 ka in the Yangtze Valley, China (Zhang et al.,
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40 2023), high foraminifer $\delta^{18}\text{O}$ values at 4.1 and 3.95 ka in the Arabian Sea indicative of weak Indian
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42 monsoon (Giesche et al., 2019), three aridity events at 4.19, 4.11, and 4.02 ka reconstructed for
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44 north India (Giesche et al., 2023), two wet anomalies from 4.16 to 4.01 ka in Namibia (Railsback et
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46 al., 2018), and anomalously wet/overcast conditions at 4.14-4.05 ka and 3.97 ka in northern Europe
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48 (Helama and Oinonen, 2019), which the latter authors called Meghalayan anomaly. Moreover, Weiss
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50 (2022) construed a series of the Indian monsoon megadroughts at 4.2-4.17 ka, 4.14-4.08 ka, and
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52 4.06-3.97 ka from the palaeoclimate literature. This set of literature shows that double-peak
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54 conditions recorded in India and China were dry, while those in northern Europe and southern Africa
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56 were wet. Notably, the second stage of the double peak would appear temporally consistent with
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3 indications of climatic disturbances specifically at 4.0 ka, as previously discussed in fewer literature
4 sources under the terms '4000 BP event' (Perry and Hsu, 2000) and '4.0 ka event' (Armitage et al.,
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6 2015).
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12 Regarding the poor age constraints minded by Parker and Harrison (2022), it is obvious that reliable
13 comparisons of the described short-term anomalies are becoming increasingly limited by
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15 uncertainties inherent to radiometric dating techniques. While uncertainties of only 10-30 years are
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17 reported in some studies (Railsback et al., 2018; Walker et al., 2018), calendar-year age ranges for
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19 calibrated ^{14}C dates around 4.2 ka typically are of order 80-150 years (Guilderson et al., 2005), for
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21 which additional uncertainties inherent to age-depth models (Telford et al., 2004) are to be added in
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23 the case of sedimentary records. Accepting the exact tree-ring dates at 4.14-4.05 ka and 3.97 ka
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25 (Helama and Oinonen, 2019) would stand as an alternative form of chronological assessment,
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27 against which the radiometric dates of other proxy records could be compared until more similarly
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29 dated high-resolution archives become available. Intriguingly, the only hitherto published tree-ring
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31 chronology from the Asian monsoon region covering the period of interest was not indicative of the
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33 4.2 ka event but instead showed a decrease in moisture availability between 4.0 and 3.5 ka (Yang et
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35 al., 2021). This would concur with Scroxton et al. (2023) who proposed a summer rainfall drying for
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37 3.97-3.4 ka around the Indian Ocean basin, following winter-season droughts of the 4.2. ka event.
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46 Late Quaternary context and global-scale forcing mechanisms
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51 Moreover, there are proxy records indicating that the change lasted longer than few centuries
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53 suggesting that the event may not have been a sudden, isolated event but developing over a much
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55 longer period. With these regards, a "tripartite" climatic development has been inferred from
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57 Mediterranean proxy records between 4.3 and 3.8 ka (Magny et al., 2009) characterized a
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59 continuum of gradual aridification, severe drought, and return to wetter conditions (Glais et al.,
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3 2023). Additional results show that the drought in India may have lasted from 5 to 3 ka, with
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5 stronger drying between 4.8 and 3.5 ka (Roy et al., 2022; Scroxton et al., 2023; Behera and Chauhan,
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7 2023). Recently, a gradual change to drier climate was reconstructed since 4.8 ka also in China (Niu
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9 et al., 2023), whereas in Ethiopia drying may have started around 5.2 ka (Lanckriet et al., 2017).
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11 Long-term change has also been inferred for pluvial conditions in Namibia since around 4.8 ka, with
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13 continuation until 2.5 ka (Sobol et al., 2022). Sensitivity of site conditions to different atmospheric
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15 patters such as ENSO variability and Inter-Tropical Convergence Zone (ITCZ) migrations has been
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17 suggested to explain the dating offsets in the Asian monsoon region (Dutt et al., 2021; Lin et al.,
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19 2022) and Africa (Lanckriet et al., 2017; Railsback et al., 2022).
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25 On the other hand, as Cruz et al. (2021) pointed out, the long-term indications predating the event
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27 highlight possibility that climate system is moving toward regimes more sensitive to an abrupt
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29 change to occur around 4.2 ka. With these regards, North Atlantic forcings factors (Bond et al., 1997,
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31 2001; Bianchi and McCave, 1999; Thornalley et al., 2009) appear the most commonly discussed
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33 factor attributed to the 4.2 ka event. This evidence represent sites not restricted to any particular
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35 region or proxy type (e.g. Sallun et al., 2012; Sun and Feng, 2013; Fubelli et al., 2013; López-Sáez et
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37 al., 2014; Walczak et al., 2015; Sarti et al., 2015; Tang et al., 2015; Blundell et al., 2018; Català et al.,
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39 2019; Tan et al., 2020). Actually, the ‘Bond 3 climate event’ and ‘Holocene event 3’ appear as
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41 occasionally used synonyms for the 4.2 ka event (Sun and Feng, 2013; Di Rita et al., 2018; Mighall et
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43 al., 2023). This forcing is commonly understood to drive a southward ITCZ shift at global scale
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45 (Broccoli et al., 2006; Yan and Liu, 2019), thus resulting in decreased rainfall over most of the Asian
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47 monsoon region (Zhang et al., 2020; Lin et al., 2022). Again, this would accord with the “cool poles,
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49 dry tropics” configuration during the 4.2 ka event (Mayewski et al., 2004). The monsoonal response
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51 to this forcing is accomplished within 110 years (Liu et al., 2018), which may partially explain the
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53 gradual long-term change in proxy records. Another reason could be that the Bond 3 event may have
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55 started to develop already around 4.6 ka (Figure 1c). It is essential to note that the North Atlantic
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3 forcing ties the 4.2 ka event to a series of similar events and climatic shifts through the Holocene
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5 with a cyclicity close to 1470 ± 500 years suggestively driven by solar forcing (Bond et al., 2001).
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7 Similar to the 4.2 ka event, the Bond 5 event at 8.2 ka (Alley and Ágústadóttir, 2005; Rohling and
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9 Pälike, 2005) and the Bond 1 event at 1.5 ka (Dark Ages Cold Period; Helama et al., 2017, 2021)
10
11 contain shorter-term anomalies that punctuate a longer and less extreme period of gradually
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13 changing climate.
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18 Also records related to ENSO (Figure 1d) and North Atlantic Oscillation (NAO) (Figure 1e) indicate
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20 small-scale perturbations between 4.2 and 4.0 ka. The NAO and the east Atlantic/west Russia
21
22 (EAWR) patterns may have affected the climatic conditions around 4.2 ka, especially in the
23
24 Mediterranean environments (e.g. Di Rita et al., 2018; Bini et al., 2019; Schirrmacher et al., 2019).
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26 However, the discussion remains superficial, not least due to deficit of NAO- and EAWR- related
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28 data. Increased ENSO variability has been recorded from around 4.4 ka towards the present (e.g. Du
29
30 et al., 2021), with a complicated period of alternating El Niño/La Niña-like conditions at interdecadal
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32 time scales between 4.1 and 3.9 ka (Dang et al., 2020). These variations in ENSO have been related
33
34 to intensification (Booth et al., 2005; Fisher et al., 2008; Li et al., 2018b) and/or temporal extension
35
36 of the anomalous climatic conditions since 4.2 ka (Scroxtton et al., 2023). According to Fisher et al.
37
38 (2008), the timing of the 4.2 ka event seems to have “inaugurated” the modern ENSO world.
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40 Combined, an interplay of NAO and ENSO in modulating the Indian summer monsoon intensities
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42 over Holocene time scales has also been proposed (Banerji and Padmalal, 2022). Volcanic forcing is
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44 not discussed as an agent behind the 4.2 ka event, which is not surprising considering the
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46 quiescence of volcanism over this period. The volcanic forcing index (Kobashi et al., 2018) obtains
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48 near-zero values during the 4.2 ka event (Figure 1f) indicating virtually no influence from this factor.
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57 Future directions
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3 Climatic events discussed under the term '4.2 ka event' represent various geographical regions and
4 types of palaeoclimate information. The sites from Mediterranean to India and China are best
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6 replicated, whereas the sites outside this region, especially from high latitudes, remain poorly
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8 represented but their records are crucially needed for more detailed assessments of the event
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10 globally. Expressing the timing of the event simply as "4.2 ka" remain too simplistic and more
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12 detailed characterization of the findings is highly recommended to allow for palaeoclimatic and
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14 chronostratigraphic assessments of the event. Some weight should be assigned in developing the
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16 reconstructions of ENSO, NAO and EAWR variability and ITCZ migrations, to validate sensitivities of
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18 the growing proxy setup to relative roles of forcing factors, and to perform proxy–model
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20 comparisons. Probably the hardest barrier will be the development of well-dated high-resolution
21
22 records to overcome the poor age constraints and to replicate the multiple proxy indications that
23
24 seemingly took place during the event, as well as to disentangle the nucleus of the event (4.2-4.0 ka)
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26 from the more gradual, long-term climatic development. Such research seems to be just beginning.
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28 Well-dated high-resolution records will not only be needed in natural science but to improve
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30 integration in societal consequences to climate change.
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39 Acknowledgements

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43 The manuscript was written while the author was supported by Grants 339788 and 355268 from the
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45 Academy of Finland.
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FIGURE CAPTIONS

Figure 1. Palaeoclimate records. Appearance of the 4.2 ka event in low-latitude and high-latitude sites of the Gulf of Oman (a) and the Mt Logan in Yukon (b) where the event is indicated by dolomite % wt. and CaCO₃ % wt. (Cullen et al., 2000) and $\delta^{18}\text{O}$ (Fisher et al., 2008) as proxies for aeolian activity and moisture, the stacked North Atlantic multi-core record of percentage of hematite-stained grains (% HSG) indicating the ice-rafted debris i.e., 'Bond events', juxtaposed with the ^{14}C production rate (pr, atoms/cm²/sec) indicating variations in solar activity, both (% HSG and ^{14}C PR) as analysed and presented in Bond et al. (2001), standardized to z-scores (Bond event 1, 3, and 5 are depicted) (c), a multi-proxy lake sediment record indicating the North Atlantic Oscillation based on lake sediments (NAO_{PCA3}) from southwestern Greenland available since 5.2 ka (Olsen et al., 2012) (d), a marine sediment record (lithic flux rate, %) off coastal Peru with El Niño–Southern Oscillation correspondence (Rein et al., 2005) (e), volcanic forcing index (VII) presented as 100-year non-overlapping means (Kobashi et al., 2017) (f), the $\delta^{18}\text{O}$ chronology from Mawmluh Cave speleothem (Berkelhammer et al., 2012) ratified as the Global Stratotype Section and Point for the lower boundary of the Meghalayan Age/Stage (Walker et al., 2018, 2019a, 2019b) (g), speleothem $\delta^{13}\text{C}$ chronology from in the Yangtze Valley, China, indicating two dry pulses (Zhang et al., 2023) (h), foraminifer $\delta^{18}\text{O}$ chronology from the Arabian Sea indicating two stages of weak Indian monsoon (Giesche et al., 2019) (i), $\delta^{18}\text{O}$ chronology of speleothem aragonite from north India shown with two dry anomalies (Giesche et al., 2023) (j), speleothem $\delta^{18}\text{O}$ chronology from Namibia indicated two-pulsed wet event depicted in low-resolution (thick line) (Sletten et al., 2013) and high-resolution data (thin line) as analysed and presented in Railsback et al. (2018) (k), and tree-ring $\delta^{13}\text{C}$ chronology (Helama et al., 2018) from northern Europe indicating two anomalies during which wet/overcast conditions prevailed (Helama and Oinonen, 2019) (l). Light and dark grey shading shown in right-hand plots indicate the timeframes between 4.4. and 3.8 ka and 4.2 and 4.0 ka, respectively. Gray

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3 shading in left-hand plots indicate the timings of two isotopic excursions recorded within the 4.2 ka
4 event. Horizontal axis: BP years (before AD 1950).
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10 Figure 2. Characterization of the proxy setup. Number of published studies reviewed in this paper
11 (a), and proportions of different geographical regions attributable to the 4.2 ka event (b) in the
12 palaeoclimate literature, with corresponding proportions of reported climatic and environmental
13 characteristics (c) and proxy types (d). Note that the number of proxy indications exceeds the
14 number of studies: in the case of multiple indications per study, they were all included, with equal
15 weight.
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25 Figure 3. Geographical distribution of different climatic and environmental proxy indications during
26 to the 4.2 ka event. The regions indicated by the pie charts are same as those given in Figure 2b. In
27 the case of multiple indications per study, they were all included, with equal weight.
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34 Figure 4. Temporal distribution of dating placements shown as (a) number of dates and as (b) bars
35 covering a window with average and median starting and ending dates attributable to corresponding
36 climatic and environmental events. The starting and ending years of the event used in these
37 calculations were adopted as reported in the original papers. Horizontal axis: BP years (before AD
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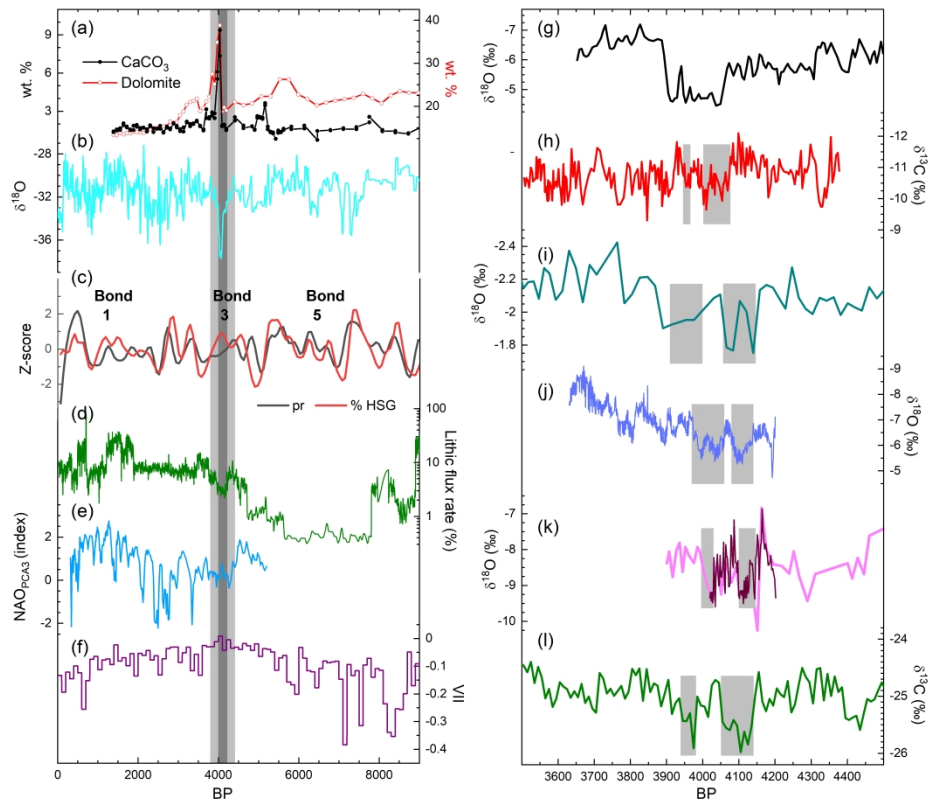


Figure 1.

280x229mm (600 x 600 DPI)

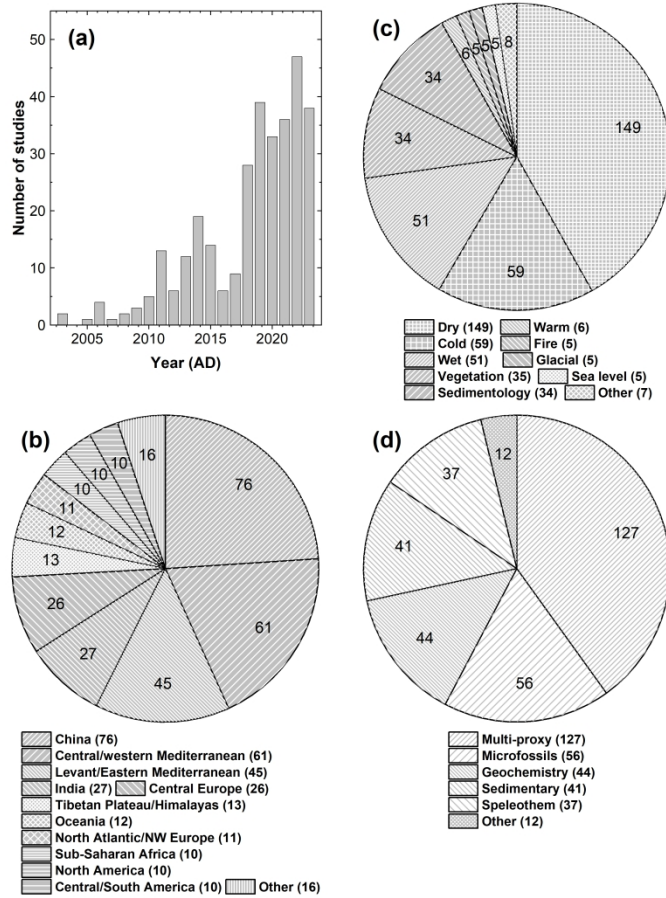


Figure 2.

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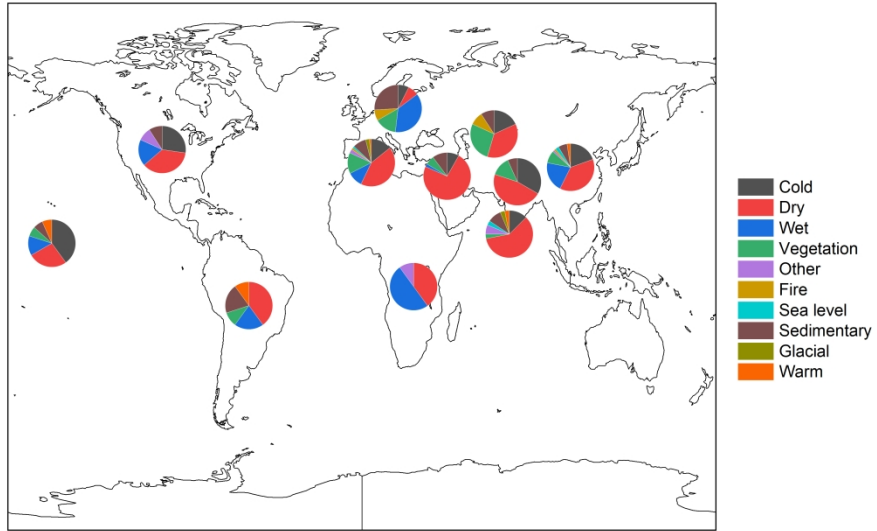


Figure 3.

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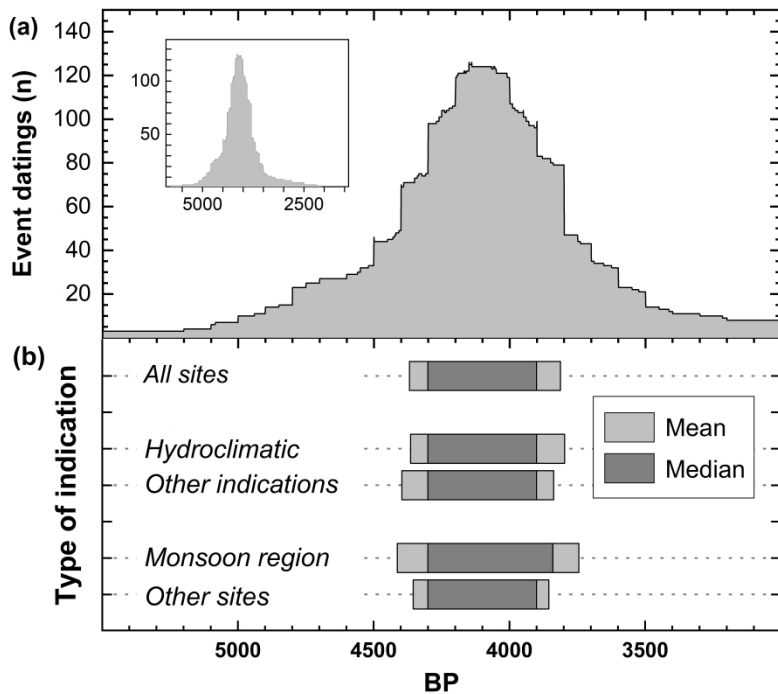


Figure 4.

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3 Supplementary Information
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5 The 4.2 ka event: a review of palaeoclimate literature and directions for future research
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8 Samuli Helama
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13 Content:
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15 Table S1. List of papers

16 Table S2. List of synonyms
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Table S1. List of papers included in the literature review.

Author(s)	Year	Title
Alexandrowicz	2023	Application of malacological analysis to reconstruct climate fluctuations and human activity during the Middle and Late Holocene. Research in the valley of the ...
Alinezhad et al.	2021	Lake Neor reveals how mountain vegetation responded to 7000 years of hydroclimate variability in northwestern Iran
Araus et al.	2014	Agronomic conditions and crop evolution in ancient Near East agriculture
Arnaud et al.	2012	Lake Bourget regional erosion patterns reconstruction reveals Holocene NW European Alps soil evolution and paleohydrology
Arz et al.	2006	A pronounced dry event recorded around 4.2 ka in brine sediments from the northern Red Sea
Avnaim-Katav et al.	2019	A multi-proxy shallow marine record for Mid-to-Late Holocene climate variability, Thera eruptions and cultural change in the Eastern Mediterranean
Azenoud et al.	2022	Climate controls on tufa deposition over the last 5000 years: A case study from Northwest Africa
Badino et al.	2018	8800 years of high-altitude vegetation and climate history at the Rutor Glacier forefield, Italian Alps. Evidence of middle Holocene timberline rise and glacier ...
Bajo et al.	2017	Stalagmite carbon isotopes and dead carbon proportion (DCP) in a near-closed-system situation: An interplay between sulphuric and carbonic acid dissolution
Baldini et al.	2019	North Iberian temperature and rainfall seasonality over the Younger Dryas and Holocene
Basavaiah et al.	2015	Late Quaternary environmental and sea level changes from Kolleru Lake, SE India: Inferences from mineral magnetic, geochemical and textural analyses
Benson et al.	2021	A speleothem record from Portugal reveals phases of increased winter precipitation in western Iberia during the Holocene
Berger et al.	2012	Rivers of the Hadramawt watershed (Yemen) during the Holocene: Clues of late functioning
Berger et al.	2013	The dynamics of mangrove ecosystems, changes in sea level and the strategies of Neolithic settlements along the coast of Oman (6000–3000 cal. BC)
Berkelhammer et al.	2012	An abrupt shift in the Indian monsoon 4000 years ago
Bernal-Wormull et al.	2023	New insights into the climate of northern Iberia during the Younger Dryas and Holocene: The Mendukilo multi-speleothem record
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Constantine et al.	2019	Mid-to late Holocene cooling events in the Korean Peninsula and their possible impact on ancient societies
Costas et al.	2014	Climate-driven episodes of dune mobilization and barrier growth along the central coast of Portugal
Courtin et al.	2021	Vegetation changes in southeastern Siberia during the Late Pleistocene and the Holocene
Coussin et al.	2023	Land-sea linkages on the Algerian Margin over the last 14 kyrs BP: Climate variability at orbital to centennial timescales
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Czerwiński et al.	2020	Late Holocene transformations of lower montane forest in the Beskid Wyspowy Mountains (Western Carpathians, Central Europe): a case study from Mount Mogielica
Dabhi et al.	2022	Mid-late Holocene climatic reconstruction from coastal dunes of the western Kachchh, India
Dang et al.	2020	El Niño/Southern Oscillation during the 4.2 ka event recorded by growth rates of corals from the North South China Sea
Das et al.	2022	Evidence for seawater retreat with advent of Meghalayan era (~ 4200 a BP) in a coastal Harappan settlement
Dash et al.	2022	An environmental magnetic record of Holocene climatic variability from the Chilika Lagoon, Southern Mahanadi Delta, east coast of India
De Falco et al.	2022	Evolution of a single incised valley related to inherited geology, sea level rise and climate changes during the Holocene (Tirso river, Sardinia, western ...
Dean et al.	2015	Eastern Mediterranean hydroclimate over the late glacial and Holocene, reconstructed from the sediments of Nar lake, central Turkey, using stable isotopes and ...
Delibes de Castro et al.	2015	The archaeological and palynological record of the Northern Plateau of Spain during the second half of the 3rd millennium Bc
Di Lorenzo et al.	2021	Human impact and landscape changes between 3000 and 1000 BC on the Tropea Promontory (Calabria, Italy)
Di Lorenzo et al.	2023	A high-resolution record of landscape changes and land use over the last 5000 years in western Calabria (S. Eufemia Gulf, southern Tyrrhenian Sea, Italy)
Di Rita & Magri	2019	The 4.2 ka event in the vegetation record of the central Mediterranean
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Gao et al.	2023	Phytolith evidence for changes in the vegetation diversity and cover of a grassland ecosystem in Northeast China since the mid-Holocene
García et al.	2022	Paleoenvironmental changes of the last 16,000 years based on diatom and geochemical stratigraphies from the varved sediment of Holzmaar (West-Eifel Volcanic ...
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Gautam et al.	2021	Indian monsoon variability during the last 46 kyr: isotopic records of planktic foraminifera from southwestern Bay of Bengal
Geirsdóttir et al.	2019	The onset of neoglaciation in Iceland and the 4.2 ka event
Ghosh et al.	2023	Last 10 millennial history of Indian summer monsoon in the Bengal region—a multi-proxy reconstruction from a lacustrine archive
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Kaal et al.	2011	Long-term deforestation in NW Spain: linking the Holocene fire history to vegetation change and human activities
Kabot-Bahr et al.	2021	A tale of shifting relations: East Asian summer and winter monsoon variability during the Holocene
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Karachurina et al.	2023	Terrestrial vegetation and lake aquatic community diversity under climate change during the mid-late Holocene in the Altai Mountains
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Kawahata	2019	Climatic reconstruction at the Sannai-Maruyama site between Bond events 4 and 3—implication for the collapse of the society at 4.2 ka event
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Stobbe & Gumnior	2021	Palaeoecology as a Tool for the Future Management of Forest Ecosystems in Hesse (Central Germany): Beech (<i>Fagus sylvatica</i> L.) versus Lime (<i>Tilia cordata</i> Mill.)
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Sun et al.	2019	The coupled evolution of mid-to late Holocene temperature and moisture in the southeast Qaidam Basin
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Susini et al.	2023	Holocene palaeoenvironmental and human settlement evolution in the southern margin of the Salpi lagoon, Tavoliere coastal plain (Apulia, Southern Italy)
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Tan et al.	2018	Centennial-to decadal-scale monsoon precipitation variations in the upper Hanjiang River region, China over the past 6650 years
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Tan et al.	2023	Levoglucosan and its isomers in terrestrial sediment as a molecular markers provide direct evidence for the low-temperature fire during the mid-Holocene in the ...
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Xie et al.	2021	Changes in the hydrodynamic intensity of Bosten Lake and its impact on early human settlement in the northeastern Tarim Basin, Arid Central Asia
Yan & Wünnemann	2014	Late Quaternary water depth changes in Hala Lake, northeastern Tibetan Plateau, derived from ostracod assemblages and sediment properties in multiple sediment ...
Yang et al.	2015	Groundwater sapping as the cause of irreversible desertification of Hunshandake Sandy Lands, Inner Mongolia, northern China
Yang et al.	2019	Total photosynthetic biomass record between 9400 and 2200 BP and its link to temperature changes at a High Arctic site near Ny-Ålesund, Svalbard
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Yang et al.	2023	Genomic investigation of the Chinese alligator reveals wild-extinct genetic diversity and genomic consequences of their continuous decline
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Ülgen et al.	2012	Climatic and environmental evolution of Lake Iznik (NW Turkey) over the last~ 4700 years
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Zanchetta et al.	2011	Tephrostratigraphy, chronology and climatic events of the Mediterranean basin during the Holocene: an overview
Zanchetta et al.	2016	The so-called “4.2 event” in the central Mediterranean and its climatic teleconnections
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Zanon et al.	2019	Palaeoenvironmental dynamics at the southern Alpine foothills between the Neolithic and the Bronze Age onset. A multi-proxy study from Bande di Cavriana (Mantua ...
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Table S2. List of synonyms for the 4.2 ka event used in the literature.

4.2 ka event
4.2 ka BP event
4.2 ka calBP event
4.2-kiloyear event
4.2-kiloyear BP event
2200 BC event
2200 BCE event
4200 BP event
4200 cal BP event
4200 calBP event
4.2 kyr event
4.2 kyr BP event
4.2 kyr cal BP event
4.2 kya event