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# Soil-tree-atmosphere CH<sub>4</sub> flux dynamics of boreal birch and spruce trees during spring leaf-out

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## Abstract

**Aims** Studies on tree CH<sub>4</sub> exchange in boreal forests regarding seasonality and role of tree canopies are rare. We aimed to quantify the contribution of boreal trees to the forest CH<sub>4</sub> budget during spring leaf-out and to reveal the role of microbes in the CH<sub>4</sub> exchange.

**Methods** Methane fluxes of downy birch and Norway spruce (*Betula pubescens* and *Picea abies*) growing on fen and upland sites were measured together with soil CH<sub>4</sub> flux, environmental variables and microbial abundances involved in the CH<sub>4</sub> cycle. Tree

CH<sub>4</sub> fluxes were studied from three stem heights and from shoots.

**Results** The trees emitted CH<sub>4</sub> with higher stem emissions detected from birch and higher shoot emissions from spruce. The stem CH<sub>4</sub> emissions from birches at the fen were high (mean 45 μg m<sup>-2</sup> h<sup>-1</sup>), decreasing with stem height. Their dynamics followed soil temperature, suggesting the emitted CH<sub>4</sub> originated from methanogenic activity, manifested in high *mcrA* gene copy numbers, in the peat soil. Methanogens were below the quantification limit in the tree tissues. Upscaled tree CH<sub>4</sub> emissions accounted for 22% of the total CH<sub>4</sub> emissions at the fen.

**Conclusions** The variation in stem CH<sub>4</sub> flux between the trees and habitats is high, and the emissions from high-emitting birches increase as the

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spring proceeds. The lack of detection of methanogens or methanotrophs in the aboveground plant tissues suggests that these microbes did not have a significant role in the observed tree-derived fluxes. The stem-emitted CH<sub>4</sub> from birches at the fen is presumably produced microbially in the soil.

**Keywords** Boreal forest · Methane flux · Methanogens · Methanotrophs · Trees · Waterlogging

## Introduction

Methane (CH<sub>4</sub>) is one of the most abundant greenhouse gases with atmospheric mixing ratio of 1.803 ppm (Hartmann et al. 2013). The quantity of atmospheric CH<sub>4</sub> has increased over 1.5-fold since preindustrial times and continues to rapidly increase (Hartmann et al. 2013). Anthropogenic sources of CH<sub>4</sub> are considered to be well-known, while cycling of CH<sub>4</sub> in terrestrial ecosystems and its natural sources remain poorly quantified (Conrad 2009; Xu et al. 2016). During the last few decades the importance of plants in CH<sub>4</sub> cycling in different ecosystems has been recognized, and much attention has been directed to revealing their role in global CH<sub>4</sub> dynamics (Carmichael et al. 2014). Numerous studies have shown that trees are capable of emitting CH<sub>4</sub> from their stems (Terazawa et al. 2007; Pangala et al. 2013, 2014, 2015, 2017; Machacova et al. 2016; Wang et al. 2016).

Most of the research on tree CH<sub>4</sub> exchange has been performed on species from tropical or temperate vegetation zones (e.g. Gauci et al. 2010; Pangala et al. 2015, 2017; Wang et al. 2016; Pitz & Magonigal 2017; Plain et al. 2019), while studies in boreal forests are still rare (Machacova et al. 2016). In general, trees growing in wetland forests have acquired more attention as they are known to be hotspots of CH<sub>4</sub> production where trees act as conduits for soil-produced CH<sub>4</sub> (Pangala et al. 2013, 2014, 2017; Jeffrey et al. 2020a; Schindler et al. 2020; Sjögersten et al. 2020; Mander et al. 2021; Moldaschl et al. 2021). Leaf-level CH<sub>4</sub> fluxes of tree canopies are much less studied, and the few studies worldwide indicate variation between small emission (Machacova et al. 2016; Pangala et al. 2017) and small uptake (Sundqvist et al. 2012; Putkinen et al. 2021), or zero fluxes (Takahashi et al. 2012), while the processes remain unknown.

The observed connections between the stem CH<sub>4</sub> fluxes and soil wetness (Rusch and Rennenberg 1998; Machacova et al. 2013, 2016; Maier et al. 2018; Barba et al. 2019a) have led researchers to set a primary hypothesis that the CH<sub>4</sub> emitted from trees is produced by methanogens in anaerobic soils and microsites, taken up by tree roots, transported into the aboveground tree tissues and released into the atmosphere (Rusch and Rennenberg 1998). In this way trees may act as pathways for the soil-produced CH<sub>4</sub> instead of the CH<sub>4</sub> being oxidized by methanotrophic microbes on its way through the aerated soil layers (Bender and Conrad 1993; Conrad 2009). Moreover, CH<sub>4</sub> emitted from trees might be produced within the tree tissues itself, by microorganisms living in plant tissues (Zeikus and Ward 1974; Covey et al. 2012; Yip et al. 2019; Putkinen et al. 2021) and/or by plant physiological processes (Keppler et al. 2006). On the other hand, initial studies indicate also a possible CH<sub>4</sub> oxidation potential in tree tissues (Jeffrey et al. 2021a, b).

Aerobic CH<sub>4</sub> production in plants has mostly been studied in laboratory conditions with herbaceous plants. Aerobic CH<sub>4</sub> emissions from herbaceous as well as woody plants have been suggested to occur especially during the growth and decay of new plant cells, linking the CH<sub>4</sub> formation to its identified CH<sub>4</sub> precursors such as lignin, cellulose, pectin and methionine (Keppler et al. 2008; Vigano et al. 2008; Messenger et al. 2009; Althoff et al. 2014; Fraser et al. 2015; Lenhart et al. 2015; Benzing et al. 2017). The importance of the processes linked to aerobic CH<sub>4</sub> production and emissions in nature are not yet fully understood.

Boreal upland forests are commonly considered a net sink for atmospheric CH<sub>4</sub> due to microbial CH<sub>4</sub> oxidation in soil, whereas the role of trees in the CH<sub>4</sub> dynamics has received relatively little attention. Boreal trees have been observed to both emit (Machacova et al. 2016; Tenhoviirta et al. 2022) and uptake (Sundqvist et al. 2012; Putkinen et al. 2021) atmospheric CH<sub>4</sub>, and these contradictory results call for more research. Furthermore, very little is known of the seasonality and the effects of tree physiological activity and environmental drivers on CH<sub>4</sub> flux dynamics of tree stems and canopies.

In this study, we measured the CH<sub>4</sub> fluxes of the stems and shoots of two common boreal tree species, downy birch (*Betula pubescens* Ehrh.) and Norway

spruce (*Picea abies* (L.) H. Karst.), as well as of the forest floor, during the spring leaf-out period. Downy birch has distribution from Scandinavia to Central Europe and to Eastern Siberia, while Norway spruce grows throughout Scandinavia and eastern parts of Europe limiting to Ural Mountains. We aim to I) evaluate the effect of the leaf-out period to the CH<sub>4</sub> exchange of the tree leaves and stems, II) assess the contribution of these tree species to the CH<sub>4</sub> cycle within a boreal forest ecosystem, and III) to connect the CH<sub>4</sub> fluxes to meteorological data, soil and tree-physiological parameters, and the abundance of the CH<sub>4</sub> producing and oxidizing microbes in the soil, within the trees and in the ground vegetation.

## Materials and methods

### Site description and experimental design

The tree CH<sub>4</sub> exchange measurements were conducted at the SMEAR II station (*Station for Measuring Forest Ecosystem-Atmosphere Relations*) in Hyttälä, southern Finland (61°51'N, 24°17'E, 181 m a.s.l.) (Hari and Kulmala 2005). Mean annual temperature and precipitation of the area are 3.5 °C and 711 mm, respectively (in 1981–2010) (Pirinen et al. 2012).

Two experimental plots differing in their forest structure and soil type and characteristics were selected within the studied forest: one situated on organic and the other on mineral soil. The organic soil site is a small (ca. 300 m<sup>2</sup>) mesotrophic forested fen with approximately 0.6 m thick peat layer and had a water table level (WTL) located 5–10 cm below the surface during the study period. Trees are unevenly aged and the main tree species in the highest canopy level is Norway spruce with downy birch as mixed

species. The hollows are dominated by *Sphagnum girgensohnii* while hummocks are covered with mix of *Pleurozium schreberi*, *Polytrichum commune* and *Hylocomium splendens* (Table 1). The upland mineral soil site is located in a Scots pine (*Pinus sylvestris*) dominated stand established in 1962 (Hari and Kulmala 2005). The soil water content at the upland site was on average 0.28 m<sup>3</sup> m<sup>-3</sup> during the study period. Downy birch and Norway spruce occur as mixed species and with Scots pine they form an evenly aged canopy layer. Understory is a sparse mix of downy birch, Norway spruce and rowan (*Sorbus aucuparia*). (The main species in the field and ground layers are listed in Table 1.)

We measured stem and shoot CH<sub>4</sub> fluxes from five trees in total: at the fen site a scaffold tower was placed so that we had access to two mature downy birches (trees number 1 and 2) and two Norway spruces (trees no. 3 and 4), and at the upland site we had access to one downy birch (tree no. 5). The shoot measurements required access to the canopy, limiting the selection of sample trees. Spruce shoots were not within range of the scaffold tower at the upland site. At the forested fen site the height of the studied trees varied between 8.5–12.9 m and the stem diameter between 0.073–0.125 m (diameter at breast height, DBH). The birch at the upland site was 19.6 m of height and 0.195 m of diameter (Table 2). Based on a forest inventory we made, the number of birch and spruce stems at the fen site were 2267 ha<sup>-1</sup> and 1400 ha<sup>-1</sup>, respectively, and the number of birches at the upland site was estimated to be 200 ha<sup>-1</sup> (including all the trees higher than 1.3 m; Table 2).

The fluxes were measured from three stem heights, and from 1–3 shoots per tree (Table 3). The forest floor CH<sub>4</sub> fluxes were measured from three points at the fen site and from one point at the upland site. The CH<sub>4</sub> flux measurement campaign was conducted in

**Table 1** Common plant species found in different parts of the ecosystem at the forested fen and upland forest sites

	Forested fen	Upland forest
Canopy	<i>Picea abies</i> , <i>Betula pubescens</i>	<i>Pinus sylvestris</i> , <i>B. pubescens</i> , <i>P. abies</i>
Understory	<i>Sorbus aucuparia</i> , <i>Salix</i> sp.	<i>B. pubescens</i> , <i>P. abies</i> , <i>S. aucuparia</i>
Field layer	<i>Equisetum sylvaticum</i> , <i>Dactylorhiza maculata</i>	<i>Vaccinium myrtillus</i> , <i>Vaccinium vitis-idaea</i> , <i>Calluna vulgaris</i>
Ground layer	<i>Sphagnum girgensohnii</i> , <i>Pleurozium schreberi</i> , <i>Polytrichum commune</i> , <i>Hylocomium splendens</i>	<i>H. splendens</i> , <i>P. schreberi</i> , <i>Cladonia</i> sp.

**Table 2** Biometric parameters of the sample trees and average trees, and the number of stems at the study sites

Plot	Species	Sample trees			Average tree per plot <sup>1</sup>		Number of trees (ha <sup>-1</sup> )
		Tree ID	Height (m)	DBH (m)	Height (m)	DBH (m)	
Forested fen	<i>B. pubescens</i> (Downy birch)	1	11.3	0.095	6.1	0.042	2267
		2	12.9	0.086			
	<i>P. abies</i> (Norway spruce)	3	8.5	0.073			
		4	12.1	0.125			
Upland forest	<i>B. pubescens</i> (Downy birch)	5	19.5	0.196	19.5	0.195	200

<sup>1</sup> Average-tree parameters are based on the site inventory, including all the trees higher than 1.3 m

the beginning of the growing season from 28 April to 11 June 2015. During the campaign, the CH<sub>4</sub> exchange measurements of the sample trees and the forest floor were conducted simultaneously twice a week at the fen and weekly at the upland forest site, except for the upland forest floor, which had semi-automated daily measurements. The measurements were mainly conducted during daytime (between 9 a.m. and 3 p.m.), complemented with five night-time measurements (between 9 p.m. and 5 a.m.) in May (on 5<sup>th</sup>, 6<sup>th</sup>, 11<sup>th</sup>, 20<sup>th</sup>, and 25<sup>th</sup> of May). All the measurements on a site were conducted within one day, and the sites were measured on successive days.

### Chamber design

The closed static chamber method (non-steady-state chambers) as described by Livingston & Hutchinson (1995) was used to determine the CH<sub>4</sub> exchange of the trees and the forest floor. Different types of stem chambers were used at the two sites. The stem chambers used at the forested fen site covered ca. 0.3 m of the stems cylindrically, enclosing a stem surface area of 0.055–0.14 m<sup>2</sup> (system volume 2.9–5.2 L; both volume and stem surface area depends on trees' DBH; modified from Machacova et al. 2016). At the upland site each stem chamber system consisted of two plastic boxes with airtight lids (Lock & Lock, Anaheim, CA, USA) interconnected with tubing (system volume 1.5 L, enclosed stem surface area in system 0.0094–0.0099 m<sup>2</sup>; Machacova et al. 2017, 2019). The shoot chambers were cylindrical in shape with FEP (fluorinated ethylene propylene) foil walls and enclosed ca. 0.3 m of the branch from the tip of the shoot (system volume 5.2 L; Machacova et al. 2016). The total leaf area in the shoot chambers at the end of the measurement campaign ranged between

0.038–0.15 m<sup>2</sup> for the birches and 0.0064–0.021 m<sup>2</sup> for the spruces. As the campaign was launched in the beginning of the growing season, we estimated the change of the leaf area and considered it in flux calculations. The air circulation in both stem and shoot chamber systems was ensured by fans and/or pumps (V1500-GAS-12 V standard vacuum pumps, flow rate 1.1 l/min, Xavitech, Sweden; NMP 850.1.2. KNDC B, flow rate 8.1 l/min, KNF Neuberger, Germany).

The forest floor CH<sub>4</sub> fluxes were measured with chambers made of aluminium or stainless steel, consisting of a permanent collar installed in the soil and an upper chamber that was closed on top of the collar. At the forested fen site there were three manual soil chambers near the sample trees (total volume 102 L, enclosed soil surface area 0.30 m<sup>2</sup>; Vainio et al. 2021). The species composition inside the chambers at the fen site consisted of *Sphagnum girgensohnii*, *P. commune*, *Carex digitata*, *Vaccinium myrtillus*, *Trientalis europaea*, *Equisetum sylvaticum*, and *Potentilla palustris*. At the upland site, one semi-automatic soil chamber measured the CH<sub>4</sub> flux once per day during 1–31 May (total volume 83 L, enclosed soil surface area 0.32 m<sup>2</sup>). The plant species composition in the chamber consisted of *V. myrtillus*, *Vaccinium vitis-idaea*, *P. commune*, *Deschampsia flexuosa*, *P. schreberi*, and *H. splendens*. The soil chambers were equipped with a fan to ensure mixing of the headspace air, and a vent-tube to minimize pressure disturbances.

### Methane flux measurements

When measuring the stem and shoot chambers, nine gas samples were taken at time intervals of ca. 1, 30, 60, 90, 120, 160, 200, 250 and 300 min after closing the chamber. Gas samples of 20 ml were

**Table 3** Shoot, stem and forest floor CH<sub>4</sub> fluxes ( $\mu\text{g m}^{-2} \text{h}^{-1}$  of leaf, stem surface, or forest floor area, respectively) expressed as mean, standard error of the mean (SEM), median, minimum and maximum, over the measured period from 28 April to 11 June 2015

Source	Plot	Species	Tree ID	Measurement height (m) <sup>a</sup>	n <sup>b</sup>	Mean CH <sub>4</sub> flux	SEM	Median flux	Min flux	Max flux	MQL index (%) <sup>c</sup>
Shoots	Forested fen	<i>B. pubescens</i>	1	8.2–10.0	30	13	11	0.45	-44	330	6.7
			2		30	63	34	4.6 *	-56	950	20
	Upland forest	<i>P. abies</i>	3	7.2–8.6	29	24	15	0.81	-5.8	370	28
			4		13	-6.6	7.3	0.88	-94	5.6	23
Stems	Upland forest	<i>B. pubescens</i>	5	14.3–15.2	18	1.4	1.1	0.52	-7.9	12	28
			1	0.5	16	51	4.8	51 *	20	89	100
	Forested fen	<i>B. pubescens</i>	1	3.6	11	0.42	0.42	0.20	-1.3	4.3	18
			2	5.1	6	0.19	0.18	0.17	-0.36	0.98	17
	Upland forest	<i>P. abies</i>	2	0.5	16	150	28	100 *	55	430	100
			3	3.6	12	1.4	0.55	0.64 *	-0.25	6.5	42
			4	5.1	9	0.17	0.09	0.11 *	-0.07	0.90	11
			5	0.4	13	0.91	0.17	0.74 *	0.23	2.2	62
Forest floor	Upland forest	<i>B. pubescens</i>	1	2.9	7	-0.24	0.21	-0.28	-1.2	0.57	14
			2	5.4	6	1.1	0.71	0.40	-0.04	4.6	33
			3	0.4	14	0.28	0.08	0.24 *	-0.22	1.1	14
			4	2.9	6	-0.03	0.21	0.10	-0.96	0.48	17
			5	5.4	6	1.0	1.1	0.17	-0.79	6.5	33
Forest floor	Forested fen	<i>B. pubescens</i>	1	0.3	3	0.30	1.1	-0.12	-1.3	2.3	0
			2	3.5	5	2.7	2.1	0.65	-1.1	11	40
			3	7.3	5	2.0	1.7	0.76	-0.97	8.7	20
Forest floor	Upland forest	<i>B. pubescens</i>	4		47	48	9.4	21 *	-3.5	210	38
			5		25	-140	6.6	-150 *	-180	-13	96

\* According to sign test, population median is significantly different from zero ( $p < 0.05$ )

<sup>a</sup> The height of the stem and shoot chambers from the ground

<sup>b</sup> Number of measurements

<sup>c</sup> The proportion of the observations that are above the method quantification limit (MQL) of the gas chromatograph

taken manually with syringes (BD Plastik™, Becton, Dickinson and Company, New Jersey, USA) and transferred into evacuated vials (12 ml, Labco Extentainer®, Labco Limited, Wales, UK). During direct sunlight, the shoot chambers were shaded with white sheets to avoid overheating of the chamber air. The headspace temperature in the shoot chambers was recorded continuously for the flux-calculation purposes with temperature sensors connected to a logger (DL2 e data logger, Delta-T Devices, Cambridge, England).

Sampling times of the soil chambers at the forested fen site were 1, 5, 15, 25, 55 and 75 min after the chamber closure, when samples of 65 ml were taken (with syringes) and 20 ml was inserted into non-evacuated vials after flushing them with the sample gas. At the upland site, the semi-automatic chamber closed automatically and injected the gas samples into evacuated vials at 1, 5, 10, 20, 30 and 50 min after closing the chamber. Chamber headspace temperature was recorded at each sampling of the soil chambers (DT-612, CEM Instruments, Shenzhen Everbest Machinery Industry Co. Ltd., Shenzhen, China). All the gas samples were stored at +5 °C in dark before analysis.

#### Analysis of the CH<sub>4</sub> concentration and the flux calculations

The CH<sub>4</sub> and CO<sub>2</sub> concentrations of the samples were analysed with a gas chromatograph (GC) (7890A, Agilent Technologies, California, USA) with a flame ionization detector (FID) and methanizer for CO<sub>2</sub> (Pihlatie et al. 2013). The method quantification limit (MQL) (Corley 2003) of the GC was estimated to be 0.10 ppm as a change in the CH<sub>4</sub> concentration, and 151 ppm for CO<sub>2</sub>.

In order to detect and omit the outliers from the CH<sub>4</sub> concentration data, we performed a robust linear regression analysis that uses iteratively reweighted least squares with a bisquare weighting function (Holland & Welsch 1977; MATLAB R2014a). The concentration points that were given a weight value below 0.9 by the robust linear model were regarded as outliers and removed from the concentration data (Vainio et al. 2021). The CH<sub>4</sub> fluxes were then calculated from the outlier-filtered data with linear fit (for the calculation, see Pihlatie et al. 2013) in relation to the tree stem surface area, the forest floor area,

and for the shoots per dry weight (DW) as well as per leaf area. The fluxes were further flagged based on the MQL, NRMSE (*Normalized Root Mean Square Error*), and R<sup>2</sup> (coefficient of determination): 1) stem and shoot fluxes above the MQL were accepted in the final data when NRMSE ≤ 0.3 and R<sup>2</sup> ≥ 0.5, otherwise they were omitted; 2) forest floor fluxes above the MQL were accepted in the final data when NRMSE ≤ 0.2 and R<sup>2</sup> ≥ 0.7, otherwise they were omitted; 3) fluxes below the MQL were accepted as such, as neither of the NRMSE or R<sup>2</sup> work for close-to-zero fluxes and omitting them would distort the data (see also Vainio et al. 2021). The forest floor CH<sub>4</sub> fluxes were further filtered based on the CO<sub>2</sub> flux data: the closures in which the CO<sub>2</sub> flux remained below the MQL, had NRMSE value higher than 0.1, or in which the CO<sub>2</sub> concentration decreased during the measurement were removed from the data. In all the stem measurements, the CO<sub>2</sub> concentration increased above the MQL. As a result, the final data comprised 88%, 93%, and 88% of the measured stem, shoot, and forest floor fluxes, respectively. In the final data, 56% of the stem fluxes, 80% of the shoot fluxes, and 42% of the forest floor fluxes were below MQL (see also Table 3).

#### Upscaling

To be able to compare the flux rates between the tree stems, shoots, and the forest floor, and in order to assess the total CH<sub>4</sub> budget of the forested fen site, the measured CH<sub>4</sub> fluxes from the trees and the forest floor were upscaled to the ecosystem level (mg CH<sub>4</sub> ha<sup>-1</sup> h<sup>-1</sup>). Upland forest was ruled out of the upscaling comparison since on the upland site we could measure only one sample tree. We first calculated the average stem surface area and crown biomass for both tree species separately, based on the height and the stem diameter of the trees at the fen site. For the average stem surface area, the tree was assumed as a cone, while the average crown biomass was calculated using equations by Repola (2008, 2009). For the spruces, the mean CH<sub>4</sub> flux of all the stem measurement heights was then upscaled for a single average-tree stem area, which was then multiplied by the number of trees per species per hectare (based on the inventory at the sites). The stem fluxes of the birches showed an exponential trend regarding the tree height (CH<sub>4</sub> emissions decreasing with height), and thus we

fitted an exponential function to the flux data. We then calculated an average flux for the birches by integrating the function from the ground level to the highest measured level, upscaled the average flux for the average-tree stem area and multiplied by the number of birches at the site.

### Environmental variables

Air temperature at 4.2 m above the ground (Pt 100), precipitation (18 m above ground, FDP12P, Vaisala, Finland), UV radiation (501A, Solar Light, USA), photosynthetically active radiation (PAR; Li-190SZ, Li-Cor, USA), soil water content in A horizon (0.02–0.06 m depth in the mineral soil, mean of five locations; TDR-100, Campbell Scientific Inc., USA), and soil temperature in A horizon (0.02–0.05 m depth in the mineral soil, mean of five locations; KTY81-110, Philips, NL) are continuously measured at the SMEAR II (available in SmartSMEAR database; Junninen et al. 2009), and these data were used as ancillary data for the upland site. For the forested fen site, soil temperature at two depths (ca. 0.05–0.10 m and 0.15–0.20 m; Thermochron iButtons, Maxim Integrated Products, USA), PAR (7–9 m above ground, Quantum sensor, Li-Cor Biosciences, USA) and sap flow of the studied trees (Granier 1987) were measured continuously during the measurement campaign. Sap flow represents an important variable describing the transpiration rate of the tree (Granier 1987; Hölttä and Kolari 2009), and was followed to investigate the relationship between the CH<sub>4</sub> fluxes and the transpiration. Soil water content was not measured at the fen site due to waterlogged conditions.

### Analyses of methanogenic and methanotrophic microbes

To evaluate the role of microbes in the CH<sub>4</sub> flux dynamics, soil, deadwood and plant samples were collected for microbial analyses from the fen and upland sites in June 2014, and at the fen site again in June 2015. The plant samples included the most abundant ground and field layer species and trees at the sites. Tree samples included separate root, stem (bore samples including layers from sap- to heartwood) and shoot tissue samples. The field layer plants were divided into below- and aboveground parts, the upland soil into litter and humus sections, and the

peat into upper (5–10 cm) and lower (30–40 cm) sections. Final samples for each soil or plant component were pooled from five subsamples. Samples were stored at –80 °C, freeze-dried and homogenized by grinding. The DNA extraction was performed from 60 mg DW (trees) or 40 mg DW (other samples) of the sample material following the procedure used for root samples by Timonen et al. (2017). Sample type origin (fen vs. upland site), and their replicate numbers (n) are listed in Tables 4 and S1.

The abundances of methanogenic archaea and methanotrophic bacteria were measured through a quantitative PCR (qPCR) analysis based on the functional genes *mcrA* (coding for the  $\alpha$ -subunit of the methyl-coenzyme M reductase) and *pmoA* (coding for the  $\alpha$ -subunit of particulate methane monooxygenase), respectively, as described in Halmeenmäki et al. (2017). The primers we used were mlas/*mcrA*-rev (Steinberg and Regan 2008) and A189f/A650r (Holmes et al. 1995; Bourne et al. 2001) for the *mcrA* and *pmoA* genes, respectively. Amplification efficiencies of the individual qPCR runs were between 91.4% and 96.8% for the *mcrA* assay, and between 86.6% and 97.7% for the *pmoA* assay.

The smallest reliably quantified standard was 10<sup>1</sup> gene copies reaction<sup>-1</sup> (both *mcrA* and *pmoA* assays) which was thus considered the limit of quantification. It should be noted that our *pmoA*-targeting approach did not cover methanotrophs lacking the *pmoA* gene, i.e. genus *Methylocella* and strains *Methyloferula stellata* AR4 and *Methyloceanibacter methanicus* R-67174 (Farhan Ul Haque et al. 2020).

### Statistical analyses

Statistical analyses were performed by using MATLAB (R2018b, MathWorks, USA). Mean, standard error of mean, minima and maxima of the CH<sub>4</sub> flux values were calculated as common statistical parameters. Normality of the distribution of CH<sub>4</sub> fluxes was tested with the Kolmogorov–Smirnov test, and due to non-normal distributions, non-parametric tests were used. The differences of medians from zero were tested with sign test. Difference between the day and night fluxes was tested with the Mann–Whitney U-test. Correlations between the CH<sub>4</sub> fluxes and the environmental variables were tested using the Spearman's rank correlation. Further analysis of correlation between the sap flow and the CH<sub>4</sub> flux was performed by comparing the sap flow to the residuals of



**Table 4** Abundances of *mcrA* and *pmoA* genes (copies g<sup>-1</sup> of sample dry weight) referring to the abundances of methanogens and methanotrophs, respectively, in the samples collected in 2014 and 2015

Sample <sup>1</sup>	Section	Site	<i>mcrA</i> (copies g <sup>-1</sup> DW)	<i>pmoA</i> (copies g <sup>-1</sup> DW)	N <sup>2</sup>
<b>2014</b>					
Peat	(5–10 cm depth)	Fen	$2.6 \times 10^7$	$5.7 \times 10^7$	1
Peat	(20–30 cm depth)	Fen	$5.1 \times 10^7$	$4.4 \times 10^7$	1
Humus soil		Upland & Fen		$4.1 \times 10^7$	1
Litter		Upland & Fen		$4.8 \times 10^6$	1
<i>Sphagnum girgensohnii</i>	Roots	Fen		$1.2 \times 10^6$	
<i>Pinus sylvestris</i>	Roots	Upland & Fen		$6.1 \times 10^5$	1
Deadwood		Upland		$4.4 \times 10^5$	1
<i>Salix</i> sp.	Roots	Upland & Fen		$3.6 \times 10^5$	1
<i>Equisetum sylvaticum</i>	Roots	Fen		$3.4 \times 10^5$	1
<b>2015</b>					
Peat	(5–10 cm depth)	Fen	$5.0 \pm 2.6 \times 10^7$	$1.6 \times 10^8 \pm 4.6 \times 10^7$	3
Peat	(20–30 cm depth)	Fen	$2.3 \pm 1.4 \times 10^8$	$4.6 \pm 2.4 \times 10^8$	3
<i>Equisetum sylvaticum</i>	Roots	Fen		$6.6 \times 10^7 \pm 7.0 \times 10^6$	3
<i>Sphagnum girgensohnii</i>	Roots	Fen		$1.1 \times 10^7 \pm 3.4 \times 10^6$	2

<sup>1</sup> Only samples, where abundances were quantifiable are shown (for unquantifiable samples, see table S1)

<sup>2</sup> The number of analysed samples (pooled from five subsamples)

fitted model between the soil temperature and stem CH<sub>4</sub> flux. Significance was assessed with a limit of  $p < 0.05$  in all statistical tests.

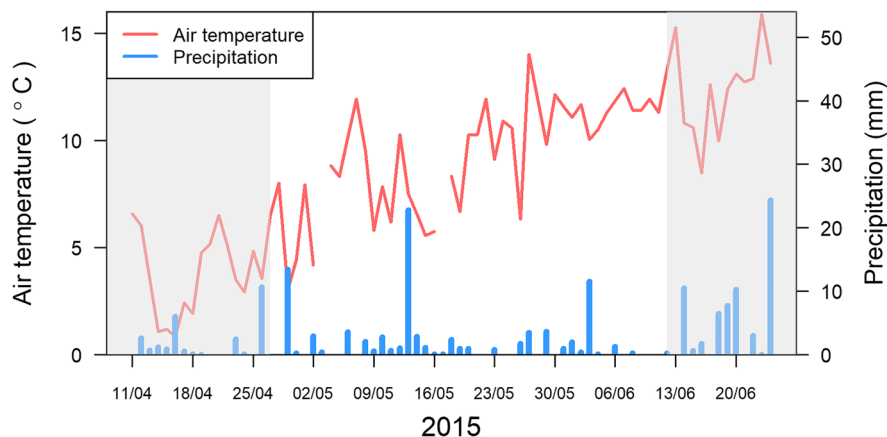
## Results

### Environmental conditions

Over the measurement campaign (from 28 April to 11 June) the daily mean air temperature had

an increasing trend from 9 to 17 May, while the increment seemed to pause in the middle of the season (Fig. 1). This is also the time when heavier precipitation/rainfall were recorded (Fig. 1). The cumulative precipitation and the mean air temperature over the measurement campaign of 8 weeks were 108.6 mm and 8.7 °C, respectively. The thermal spring turned into thermal summer on 25.5.2015 (Finnish Meteorological Institute (FMI) n.d.).

**Fig. 1** Daily averages of air temperature and precipitation, and 10 day cumulative precipitation measured at the SMEAR II station in Hyytiälä during April–June 2015. The CH<sub>4</sub> flux measurements were conducted from 28 April to 11 June 2015, which is between the grey-shadings in the figure

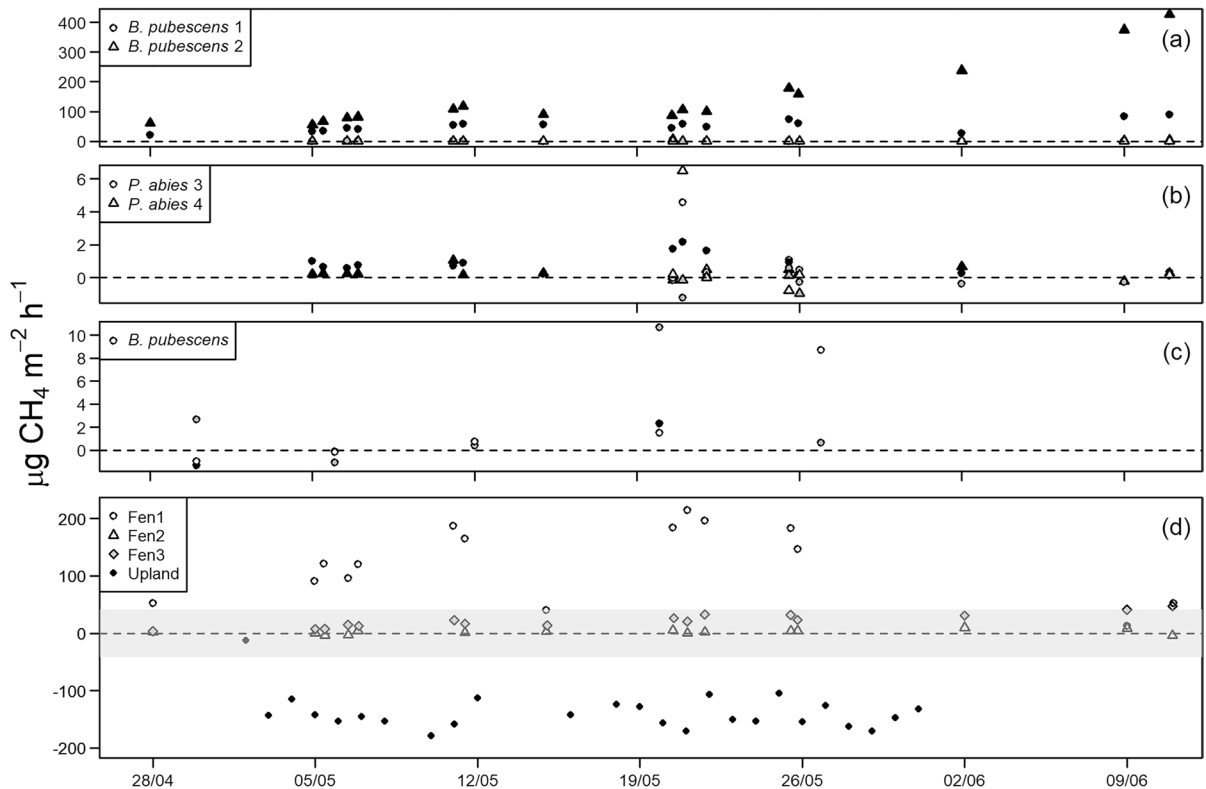


## Methane exchange of stems, shoots and forest floor

During the measurement campaign we observed both emissions and occasional uptake of  $\text{CH}_4$  by the stems and shoots of birch and spruce. The mean and median fluxes calculated over the measurement period indicated mainly emissions from the stems and shoots of both species at both sites. At the forested fen, the  $\text{CH}_4$  fluxes of the birch and spruce stems ranged from  $-1.3$  to  $430$  and from  $-1.2$  to  $6.5 \mu\text{g m}^{-2} \text{h}^{-1}$  (of stem surface area), respectively, and the median  $\text{CH}_4$  emissions from the birch stems ( $0.11$ – $100 \mu\text{g m}^{-2} \text{h}^{-1}$ , depending on the height) were mainly higher than from spruces ( $-0.28$ – $0.74 \mu\text{g m}^{-2} \text{h}^{-1}$ ) (Table 3). Regarding the birches at the fen, the emissions from the height of  $0.5$  m were substantially higher (medians  $51$  and  $100 \mu\text{g m}^{-2} \text{h}^{-1}$ , for birches n. 1 and 2

respectively; Table 3) than from the upper parts of the stems (medians  $0.11$ – $0.64 \mu\text{g m}^{-2} \text{h}^{-1}$ ; Table 3), and increased as the growing season proceeded (Fig. 2a). The spruce stems did not show such clear pattern regarding stem height (Table 3) nor temporal variation (Fig. 2b). At the stem height of  $2.9$  m,  $\text{CH}_4$  uptake by spruce dominated  $\text{CH}_4$  emission (Table 3). In contrast to the birches at the forested fen site, the stem  $\text{CH}_4$  fluxes of the birch tree at the upland site varied from  $-1.3$  to  $11 \mu\text{g m}^{-2} \text{h}^{-1}$  (Table 3), and the fluxes did not show clear temporal or height-related patterns (Fig. 2c)

The  $\text{CH}_4$  fluxes from the birch shoots at the forested fen indicated mean emissions of  $13$  and  $63 \mu\text{g m}^{-2} \text{h}^{-1}$  of leaf area from birches 1 and 2, respectively, while the mean shoot fluxes of two spruces showed both emission and uptake ( $24$



**Fig. 2** Stem  $\text{CH}_4$  fluxes ( $\mu\text{g CH}_4 \text{ m}^{-2} \text{h}^{-1}$ ) of (a) two downy birches (*B. pubescens*) and (b) two Norway spruces (*P. abies*) growing at the forested fen site, and (c) one downy birch (*B. pubescens*) growing at the upland forest site, and soil  $\text{CH}_4$  fluxes (d) at forested fen and upland site. All fluxes were measured in April–June 2015. The symbols in stem flux figures (a,b,c) are highlighted black, grey and white denoting the

measurement height of  $0.3$ – $0.5$  m,  $2.9$ – $3.6$  m and  $5.9$ – $7.3$  m of the tree stems, respectively. In soil  $\text{CH}_4$  flux figure (d), Fen1, Fen2, and Fen3 denote manual chambers at the forested fen site, and Upland denotes semi-automatic chamber at the upland site. Flux values inside the grey area were below the method quantification limit. Notice the differing y-axis scales

and  $-6.6 \mu\text{g m}^{-2} \text{h}^{-1}$  from spruces 3 and 4, respectively) (Table 3). The median flux of the fen birch no. 2 ( $4.6 \mu\text{g m}^{-2} \text{h}^{-1}$ ) was significantly different from zero ( $p < 0.05$ ), while the other trees indicated small median emissions ( $0.45\text{--}0.88 \mu\text{g m}^{-2} \text{h}^{-1}$ ). The shoot  $\text{CH}_4$  fluxes did not indicate any seasonal dynamics related to leaf growth.

We found a positive correlation between the soil temperature and the  $\text{CH}_4$  flux from the near-ground stem parts of the birch trees at the forested fen site ( $r_s = 0.61$ ,  $p < 0.05$ , and  $r_s = 0.92$ ,  $p < 0.0001$ , for the two birches separately; soil temperature at  $0.15\text{--}0.20 \text{ m}$ ; Fig. 3). The correlations were only slightly weaker with the soil temperature closer to the surface ( $r_s = 0.52$ ,  $p < 0.05$ ;  $r_s = 0.82$ ,  $p < 0.001$ ; data not shown). Otherwise, tree stem fluxes or shoot fluxes did not correlate significantly with soil temperature or with soil water content, soil  $\text{CH}_4$  fluxes, sap flow, air temperature, or radiation ( $p > 0.05$ ). We did not observe any difference between the medians of the daytime and the night-time  $\text{CH}_4$  fluxes, neither for the stems or shoots, nor the forest floor (night-time medians:  $0.62 \mu\text{g m}^{-2} \text{h}^{-1}$  (stems),  $0.0025 \mu\text{g m}^{-2} \text{h}^{-1}$  (shoots),  $110 \mu\text{g m}^{-2} \text{h}^{-1}$  (soil)).

The forest floor at the forested fen site was a net source of  $\text{CH}_4$  with a range from  $-3.5 \mu\text{g m}^{-2} \text{h}^{-1}$  to  $210 \mu\text{g m}^{-2} \text{h}^{-1}$ , and a mean of  $48 \mu\text{g m}^{-2} \text{h}^{-1}$  (Table 3). In contrast, the forest floor fluxes measured at the upland site showed  $\text{CH}_4$  uptake varying from  $-180$  to  $-13 \mu\text{g m}^{-2} \text{h}^{-1}$  with a mean value of  $-140 \mu\text{g m}^{-2} \text{h}^{-1}$  (Table 3). Only one of the three soil chambers at the forested fen site showed seasonal

increase in the  $\text{CH}_4$  emissions from April to May, while in other locations at the forested fen and at the upland site the forest floor fluxes remained rather constant (Fig. 2d).

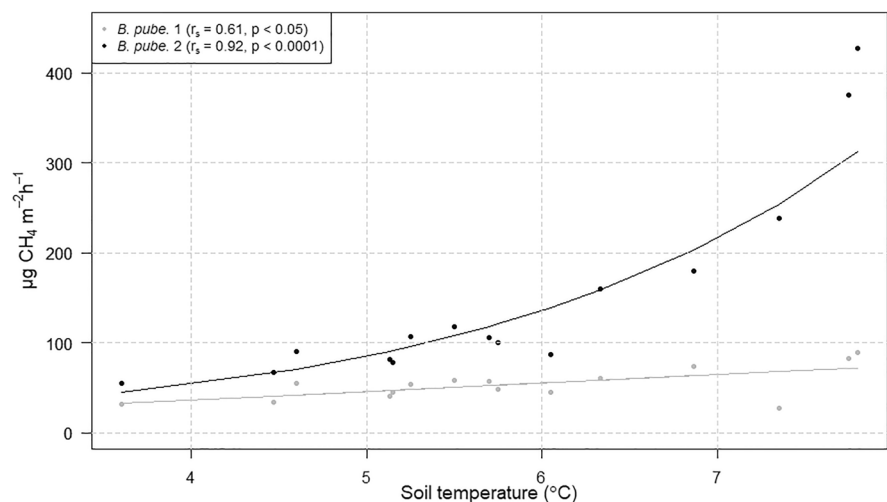
#### Upscaling of the stem and shoot fluxes at the fen site

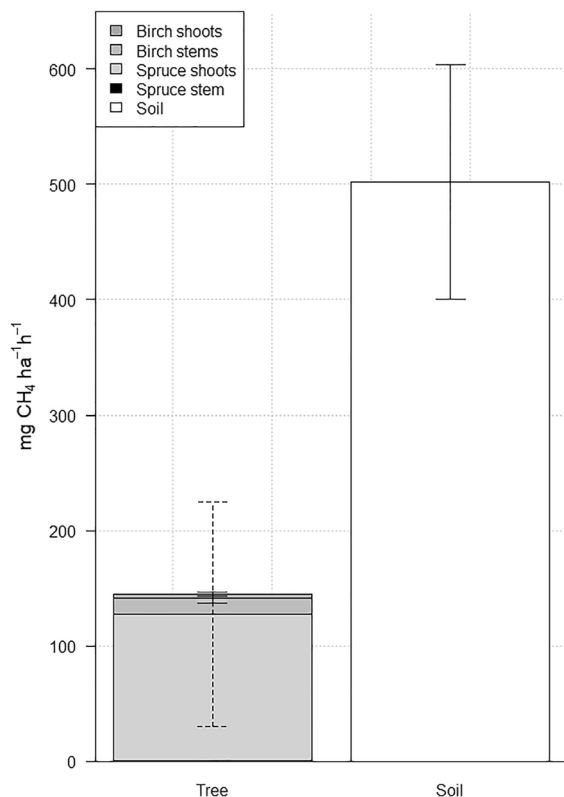
The upscaled  $\text{CH}_4$  fluxes show that the studied trees were net emitters of  $\text{CH}_4$ . The upscaled flux values (mean  $\pm$  SEM) for the stems of birch and spruce were  $14 \pm 4.0$  and  $0.62 \pm 0.20 \text{ mg ha}^{-1} \text{h}^{-1}$ , and for the shoots of birch and spruce  $3.5 \pm 1.7$  and  $127 \pm 97 \text{ mg ha}^{-1} \text{h}^{-1}$ , respectively (Fig. 4). The sum of stem and canopy  $\text{CH}_4$  emissions were  $17 \pm 5.7 \text{ mg ha}^{-1} \text{h}^{-1}$  from birch trees and  $128 \pm 97 \text{ mg ha}^{-1} \text{h}^{-1}$  from spruce trees, respectively. The forest floor emitted  $500 \pm 100 \text{ mg ha}^{-1} \text{h}^{-1} \text{CH}_4$ . The combined  $\text{CH}_4$  emission rate of the trees and forest floor was  $650 \pm 200 \text{ mg ha}^{-1} \text{h}^{-1}$ , where trees represent 22% of the total flux.

#### Abundances of methanogenic and methanotrophic microbes

The quantification of the methanogenic *mcrA* genes was successful (abundance above the quantification limit =  $10^1$  gene copies reaction $^{-1}$ ) only for the peat samples of the forested fen site, where the gene abundances varied between  $2.6 \times 10^7$  and  $2.3 \times 10^8$  copies  $\text{g}^{-1}$  (DW) within the two sampled years (Table 4). For the methanotrophic *pmoA* genes, the quantification was successful at the forested fen site for the

**Fig. 3** Linear (*B. pubescens* 1) and exponential (*B. pubescens* 2) relationships and the Spearman's correlation coefficients ( $r_s$ ) between the stem base  $\text{CH}_4$  fluxes of the birches (0.5 m above the ground) and soil temperature (in 15–20 cm depth) at the forested fen





**Fig. 4** Upscaled CH<sub>4</sub> fluxes (mg CH<sub>4</sub> ha<sup>-1</sup> h<sup>-1</sup> ground area) from stems and shoots of birch (*B. pubescens*) and spruce (*P. abies*) (left), and forest floor (right) at the forested fen (upscaled CH<sub>4</sub> fluxes of the stems of spruces are relatively small and thus not clearly visible)

peat samples (gene abundances between  $3.7 \times 10^7$  and  $4.6 \times 10^8$  copies g<sup>-1</sup> DW), and for the roots of *E. sylvaticum*, *S. girgensohnii* and *Salix* sp., which had gene abundances between  $2.3 \times 10^5$  and  $6.6 \times 10^7$  copies g<sup>-1</sup> (DW). At the upland site, *pmoA* abundances were quantifiable for the humus and litter samples, and for the *P. sylvestris* roots (between  $4.4 \times 10^5$  and  $3.4 \times 10^7$  copies g<sup>-1</sup> DW). In addition, one of the two deadwood samples collected from the upland site had a *pmoA* abundance of  $3.8 \times 10^5$  copies g<sup>-1</sup> (DW). The values represent mean abundances, while the standard errors are given in Table 4 when applicable (n=1 for 2014 and n=2 to 3 for 2015). None of the above sample types differed significantly from each other (only applicable for the 2015 samples). The analysed samples, in which these genes could not be quantified, are listed in Supporting Information (Table S1).

## Discussion

We studied the CH<sub>4</sub> exchange of common boreal trees and the effect of spring leaf-out to the fluxes. This study complements the scarce research on boreal tree CH<sub>4</sub> exchange, and especially globally rare shoot flux measurements (Machacova et al. 2016). Surprisingly, we did not observe any effect of spring recovery or leaf-out to the shoot CH<sub>4</sub> exchange. Furthermore, there was no clear temporal pattern in the stem CH<sub>4</sub> exchange, except for the bottom parts of the birch stems at the fen site. In addition, the CH<sub>4</sub> exchange of neither stems nor shoots correlated with the sap flow.

### Highly variable stem CH<sub>4</sub> emissions

In this study, measured trees showed highly variable stem CH<sub>4</sub> dynamics both in time between the measurement days and between tree species. The emissions from birch stems at the fen site followed the pattern of decreasing CH<sub>4</sub> emissions with stem height (Barba et al. 2019a), similar to earlier studies in temperate wetland and upland forests (Pangala et al. 2015; Wang et al. 2016; Jeffrey et al. 2020b; Schindler et al. 2020). The spruces at the fen and the birch at the upland site also emitted CH<sub>4</sub> from their stems, but in much smaller amounts and without a similar height-related pattern than the fen birches.

Overall, the CH<sub>4</sub> emissions from the tree stems in this study (medians between  $-0.28$ – $100$  μg m<sup>-2</sup> h<sup>-1</sup>, depending on the species, height and site) were rather small, yet in the same magnitude, compared to the flux rates reported from the tree stems in temperate zones (Terazawa et al. 2007, 2015; Pangala et al. 2015; Pitz and Megonigal 2017; Maier et al. 2018; Pitz et al. 2018; Barba et al. 2019b; Moldaschl et al. 2021). Pangala et al. (2015) reported an average summertime CH<sub>4</sub> emission of  $203 \pm 21$  μg m<sup>-2</sup> h<sup>-1</sup> from downy birch stems. In our study, the measurement period was limited to the spring and early summer, and thus we were not able to see if the emissions of birches continued to increase during the summer. Some previous studies reported substantial variety in the stem CH<sub>4</sub> exchange between tree individuals (Terazawa et al. 2015; Maier et al. 2018; Moldaschl et al. 2021; Köhn et al. 2021), similarly as the fen birches in this study.

## Sources of the stem-emitted CH<sub>4</sub>

Stem-emitted CH<sub>4</sub> has been proposed to originate either (1) from microbial production in the soil, where CH<sub>4</sub> is then taken up by tree roots, transported to the aboveground tissues and released to the atmosphere (Rusch and Rennenberg 1998), or (2) from in situ CH<sub>4</sub> production by microorganisms inside the tree stems (Zeikus and Ward 1974; Covey et al. 2012; Yip et al. 2019; Li et al. 2020). In our study, the CH<sub>4</sub> emissions from the birch stems followed soil temperature, whereas no such correlation was found with the spruce trees. Similarly, Pangala et al. (2015) reported stem CH<sub>4</sub> emissions of downy birch increasing from April to June. Several other studies have reported similar relation between the stem fluxes and soil temperature, suggesting that the CH<sub>4</sub> originates from methanogenic activity within the soil (Pangala et al. 2013, 2015; Terazawa et al. 2015). We found high abundances of both methanogens and methanotrophs in the peat profile at the fen site, demonstrating potential CH<sub>4</sub> production and oxidation in the peat profile. Forest floor CH<sub>4</sub> fluxes were close to zero in two out of three soil chambers (Fig. 2d), presumably also linked to methanotrophic activity at the peat surface (Putkinen et al. 2012). Interestingly, simultaneous CH<sub>4</sub> fluxes from especially birch stems were high indicating that birches act as conduits for the soil-CH<sub>4</sub>, hence bypassing the peat profile where CH<sub>4</sub> could be oxidized by methanotrophs. The forest floor CH<sub>4</sub> fluxes at the forested fen site were in line with previous reports from pristine spruce mires (Huttunen et al. 2003; Koskinen et al. 2016).

Pangala et al. (2014) and Maier et al. (2018) found that the amount of CH<sub>4</sub> dissolved in soil pore water was controlling the rate of stem CH<sub>4</sub> flux of black alder (*Alnus glutinosa*) saplings and mature European beech (*Fagus sylvatica*), whereas Machacova et al. (2016) found significant correlations between CH<sub>4</sub> fluxes of soil and tree stems of mature Scots pine. These studies suggest that soil CH<sub>4</sub> concentration is a critically important driver of tree stem CH<sub>4</sub> emission, however, none of the studies confirmed the CH<sub>4</sub> transport mechanism. The soil-borne CH<sub>4</sub> is suggested to be transported into the atmosphere either passively by diffusion via aerenchymatic structures and/or intercellular spaces (Machacova et al. 2013; Pangala et al. 2014; Terazawa et al. 2015; Maier et al. 2018; Plain

& Epron 2021), or actively via transpiration stream of the trees (Pangala et al. 2015; Machacova et al. 2016).

Even though there was no correlation between the stem CH<sub>4</sub> exchange and the sap flow nor between stem and soil CH<sub>4</sub> fluxes in our study, this does not rule out the possibility that CH<sub>4</sub> is transported in stems via transpiration stream. Recent modelling study by Anttila et al. (*submitted manuscript*) states that the transpiration stream is in fact the major transport pathway of axial CH<sub>4</sub> movement in tree stems, however, due to the long transit time of CH<sub>4</sub> within the stem and the resulting CH<sub>4</sub> storage in the stems, the temporal dynamics of sap flow are not visible in the stem CH<sub>4</sub> fluxes. Further, higher sap flow rate in birches (Gartner et al. 2009) could result a greater CH<sub>4</sub> storage within birch stems and partly explain higher stem CH<sub>4</sub> fluxes from the birch stems. Similar to our study, Schindler et al. (2021) did not detect any clear differences in stem CH<sub>4</sub> emissions of *Alnus incana* between the daytime and night-time. However, diurnal measurement of CH<sub>4</sub> fluxes from stems and shoots are scarce and should be considered in future work as omitting them might lead to over- or underestimation of upscaled CH<sub>4</sub> fluxes.

All the sample trees at the forested fen site are growing under similar conditions under high water table, and thus the differences in CH<sub>4</sub> exchange between the studied species may be due to different in situ CH<sub>4</sub> production mechanisms, or differences in the physiology and/or anatomy of the species. One such difference is the root development: the roots of downy birch at peatland sites grow both in vertical and horizontal directions (Huikari 1959), whereas Norway spruce is traditionally assumed to grow roots closer the soil surface (e.g. Puhe 2003; Konôpka et al. 2010). Thus, possibly only the birch roots were penetrating into the deeper soil layers that are conducive to CH<sub>4</sub> production. The angiosperms (including broad-leaved trees) are noted to cope better with waterlogging than gymnosperms (including conifers) (Kozłowski and Pallardy 1997). The primary adaptation of plants to flooding is the capacity to transport O<sub>2</sub> from the atmosphere into the roots (Kozłowski 1997) through structures such as hypertrophied lenticels, aerenchymatic tissue in the roots and/or adventitious roots (Hook 1984). It is, however, yet to be confirmed whether this pathway is potential either for downy birch or Norway spruce.

While CH<sub>4</sub> emissions at the fen site, especially by the birches, were presumably caused by the transportation of soil-derived CH<sub>4</sub>, the origin of the emitted CH<sub>4</sub> from the birch at the upland site remains unclear. It is possible that methanogens inhabit upland soils (Angel et al. 2012; Lyu and Lu 2018) and we have earlier (in 2012) measured *mcrA* gene copies between  $1.1 \times 10^5$  and  $3.6 \times 10^6$  copies g<sup>-1</sup> (DW) at moist regimes of the upland site (unpublished results) to which roots can have access. Recently, Barba and colleagues (Barba et al. 2021) reported elevated CH<sub>4</sub> concentrations and significant CH<sub>4</sub> emissions from upland trees growing on soils with below ambient CH<sub>4</sub> concentrations. This supports earlier findings that even upland trees may have CH<sub>4</sub> production within living stems, and this CH<sub>4</sub> production has been connected to methanogenic archaea living in wet and/or rotten heartwood (Yip et al. 2019; Li et al. 2020). While we could not sample the heartwood of the same trees from which we measured the fluxes, the lack of quantifiable *mcrA* genes in the comparable tree samples (Table 4 and S1) suggests that stem-inhabiting methanogens are a negligible CH<sub>4</sub> source in this boreal forest. However, we cannot completely rule out the effect of other relevant microbes not covered by the *mcrA* gene-based approach, or the presence of populations below the quantification limits (Putkinen et al. 2021).

#### Canopy CH<sub>4</sub> emissions affect the whole-tree CH<sub>4</sub> budget

All the studied tree species emitted small amounts of CH<sub>4</sub> from their shoots, but also occasional CH<sub>4</sub> uptake was observed. These findings are in line with previously reported canopy CH<sub>4</sub> exchange rates from boreal trees in Finland (Machacova et al. 2016) and in Sweden (Sundqvist et al. 2012; Putkinen et al. 2021). When the net emissions on the forested fen site were upscaled per canopy biomass of the trees, the total canopy emissions were nine-fold compared to tree stem emissions in trees, due to high leaf biomass. This indicates that even small emissions at the leaf-level may have a substantial contribution to the whole-tree CH<sub>4</sub> dynamics, and to the ecosystem CH<sub>4</sub> budget during the spring leaf-out period. As the measured shoot fluxes in both birch and spruce canopies were very small and highly variable, our findings should be considered as indicative. The used

measurement technique was able to detect CH<sub>4</sub> fluxes even with a relatively long chamber closure time, which certainly affected the physiological functioning of the tree. Excessive heating of the chamber was avoided by physical shading of the chambers, which blocked the incoming UV- and photosynthetically active radiation (PAR) to the leaves. As a result, the possible light-driven CH<sub>4</sub> emissions, found in laboratory studies (Keppler et al. 2006; Vigano et al. 2008; Fraser et al. 2015; Martel and Qaderi 2019), were likely missing. This could also explain why we did not find a connection between the shoot CH<sub>4</sub> emissions and radiation (PAR, UV). The high variability in CH<sub>4</sub> exchange by the shoots also suggests that both uptake and emission may be present as reported also by Sundqvist et al. (2012) and Putkinen et al. (2021).

The leaf-level CH<sub>4</sub> consumption detected in the previous field studies (Sundqvist et al. 2012; Putkinen et al. 2021) was hypothesized to result from methanotrophic activity in the leaves, while the presence of currently recognized methanotrophs was not confirmed in their studied trees. In our study, we were unable to detect quantifiable amount of methanotrophs in the analysed shoot samples. Still, lack of quantification does not rule out small populations, which is supported by previous detection of variable alphaproteobacterial methanotrophs in conifer needles by cultivation (Doronina et al. 2004; Iguchi et al. 2012) and 16S rRNA sequencing methods (Rúa et al. 2016; Haas et al. 2018). In the seasonal follow-up by Haas et al. (2018), needle methanotroph abundance increased towards the end of the growing season – a pattern, which could partly explain our results from the early summer. In addition, the methanotroph detected by Haas et al. (2018) was of the genus *Methylocella*, which lacks the functional *pmoA* gene targeted in our study.

Upscaling the forest floor, tree stem and canopy fluxes to the whole-forest exchange at the forested fen, we found that the trees contributed significantly to the net CH<sub>4</sub> balance. Inclusion of the CH<sub>4</sub> emissions from the tree stems and shoots increased the total CH<sub>4</sub> emissions by 29% compared to the forest floor flux of the site. Similar whole-forest upscaling, including stem and canopy CH<sub>4</sub> exchange, has only been conducted by Machacova et al. (2016) for Scots pine trees at an adjacent experimental upland forest plot at Hyytiälä. In that upscaling, the CH<sub>4</sub> emissions from Scots pine trees offset 0.8% of the CH<sub>4</sub>

sink of the upland soil. Our results indicate that at the forested fen, where the soil is a significant source of CH<sub>4</sub>, a large part of the ecosystem CH<sub>4</sub> emissions could still originate from trees. The canopy-scale CH<sub>4</sub> emissions, however, should be considered with caution due to the high uncertainties in the measured fluxes as well as in the methods to upscale to the ecosystem level. Furthermore, the number of studied trees was very limited. The upscaling here needs to be considered as one of the first efforts to understand the role of tree canopies and stems to the whole-forest CH<sub>4</sub> balance, and more measurements with higher temporal coverage and with an improved canopy flux measurement method are required for reliable estimations.

According to this study, trees at the boreal vegetation zone are capable of emitting CH<sub>4</sub> from their stems and canopies, with the emission rates in the same magnitude as those from temperate forests. The variation in stem CH<sub>4</sub> flux between tree species, tree individuals and different habitats is high, and the CH<sub>4</sub> emissions from high-emitting trees – birches on peat soil – increase as the growing season proceeds. The lack of detection of methanogens or methanotrophs in the aboveground plant tissues suggests that these microbes did not have a significant role in the observed tree-derived CH<sub>4</sub> fluxes. At least at the fen site, the stem-emitted CH<sub>4</sub> from birch trees is most likely produced microbially in the soil. More research is needed to study the possible differences between tree species. Long-term and continuous CH<sub>4</sub> flux measurements of tree stems and shoots are critically needed to accurately identify the drivers behind the CH<sub>4</sub> dynamics and to assess the contribution of trees on forest CH<sub>4</sub> budget. Finally, studies regarding trees part taking in CH<sub>4</sub> dynamics are critically important for creating a comprehensive baseline data to be used in research of climate change and its drivers.

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#### Declarations

**Conflict of interest** The authors do not have financial or non-financial conflicting interests related to the work submitted for publication.

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